Wave overtopping on a low-crested seawall under extreme waves

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9 Abstract

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10 Extreme waves in global nearshore regions, frequently accompanied by wave setup, can transform 11 seawalls into low-crested structures. Such events pose threats to coastal infrastructure due to enhanced 12 overtopping and intense hydrodynamic loads imposed on seawalls and related coastal defenses. This 13 study investigates the wave overtopping dynamics on a low-crested seawall under extreme wave 14 conditions through controlled wave flume experiments that have comprehensively measured wave 15 elevations, free-surface profiles, overtopping volumes and impact pressures. The temporal and spatial 16 characteristics of wave overtopping and their dependence on water depth and wave parameters are 17 examined. The results demonstrate a positive correlation between overtopping volume and wave 18 amplitude, with localized impact pressures also intensifying as wave amplitude increases. Conversely, 19 as wave peak frequency increases and seawall crest elevation rises, waves, especially those of larger 20 amplitudes, tend to break earlier on the seaward slope. This earlier breaking dissipates a significant 21 portion of wave energy, thereby reducing overtopping volumes and the impact pressures on the seawall. 22 Furthermore, the study reveals that as the focusing position of the extreme wave group shifts landward, 23 there is a notable reduction in the group's cumulative energy. This energy attenuation results in 24 diminished overtopping volumes and lower impact loads. These findings elucidate the complex 25 interplay between wave parameters, seawall height and the dynamics of wave overtopping under extreme wave conditions, as well as provide a theoretical framework for optimizing the design and 26 27 resilience of seawalls to mitigate the adverse impacts of extreme wave events on coastal infrastructure. 28

29 Keywords: Wave overtopping; Seawall; Wave flume experiment; Extreme wave; Wave impingement

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1. Introduction

enhancing disaster mitigation efforts.

46 Thornton (2016) analyzed overtopping volumes' temporal variation and identified a peak overtopping 47 volume several times greater than the average followed by a gradual decline. Some studies, e.g., Van 48 der Meer et al. (2010), Chen et al. (2015), Mares-Nasarre et al. (2019) and van Bergeijk et al. (2019), 49 investigated overtopping flow depths and velocities on seawalls' horizontal crests and found that the 50 overtopping flow depths decrease exponentially with distance from the crest edge although different 51 decay rates were used across studies. Gallach-Sánchez et al. (2021) conducted experimental investigations into the overtopping of steep, low-crested seawall structures induced by non-breaking 52 53 irregular waves, with an emphasis on how wave characteristics influence the distribution of overtopped 54 water and associated impact dynamics. Esteban et al. (2022) explored overtopping discharge around a 55 fixed vertical cylinder under non-impulsive wave conditions using a combination of experimental and 56 numerical methodologies. More recently, Wong and Chow (2024) performed numerical simulations 57 using OpenFOAM to analyze the wave run-up and overtopping characteristics of ocean swells 58 propagating over varying seabed bathymetries. Rif'atin et al. (2024) investigated the effectiveness of 59 stepped revetments in reducing wave run-up height and overtopping discharge by invoking the Genetic 60 Algorithm for optimization. 61 Recent studies have increasingly focused on the wave overtopping induced by a solitary wave,

62 which is a typical example of extreme wave events. For example, Hsiao and Lin (2010) investigated the 63 solitary wave overtopping and impinging process through experiments and numerical simulations and 64 highlighted the high risks of extreme wave impacts in the condition of rising water levels. Baldock et 65 al. (2012) experimentally studied the overtopping by solitary waves and found that overtopping rates

With global climate change, the frequency and intensity of extreme waves along the coast of the

world are increasing (Dysthe et al., 2008, Li et al., 2024, Lobeto et al., 2024). These extreme waves are

often accompanied by rapid coastal water level rises, transforming existing seawalls into low-crested

structures, with a relative crest elevation of $0 \le R_c/H_0 \le 1.5$ (Van der Meer et al., 2018). Such conditions

lead to excessive overtopping volumes as seawater surges over seawalls, resulting in coastal flooding,

economic losses, and structural damage (Nikolkina and Didenkulova, 2012). Wave breaking during interactions between extreme waves and seawalls generates intense slamming pressures that further

impact seawalls and nearby infrastructure (Qu et al., 2022, Wang et al., 2025). Consequently,

investigating the hydrodynamic mechanisms of overtopping on low-crested seawalls under extreme

wave conditions is essential for designing resilient seawall structures, protecting coastal cities and

research efforts have been devoted to the characteristics and prediction of overtopping volume (Paape,

1960, Owen and Steele, 1993, Van der Meer et al., 2018). Over the past two decades, advancements in

experimental methods have enhanced the understanding of seawall overtopping dynamics. Hughes and

Wave overtopping can cause severe flooding and waterlogging behind seawalls. Substantial

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increased linearly with the deficit in the wave run-up freeboard. Luo et al. (2019) have applied the 66 67 Consistent Particle Method (CPM) to simulate the solitary wave overtopping process and the results 68 have suggested that higher water levels can result in much more overtopping volume. Despite these 69 progresses, most relevant studies have focused on overtopping volumes, e.g., Goda (2009), Nørgaard 70 et al. (2014), Van Doorslaer et al. (2015), Pan et al. (2015), Hughes and Thornton (2016), Molines et 71 al. (2019), Van der Meer et al. (2018), Salauddin and Pearson (2019), etc. Some studies have been 72 devoted to examining the hydrodynamic processes of wave overtopping, e.g. nonlinear wave breaking, 73 energy dissipation along the seawall, and violent wave impacts on the seawall. However, relevant 74 studies have primarily considered regular (Wen et al., 2019, Adibhusana et al., 2023), irregular (Liu et 75 al., 2020, Koosheh et al., 2024), solitary waves (Hsiao et al., 2008, Huang et al., 2022) or tidal bores 76 (Qu et al., 2024) and few have touched extreme waves. Considering the destructive effects of extreme 77 waves on coastal infrastructures, it is crucially important to reveal the dynamic process and overtopping 78 mechanisms of extreme waves on low-crested seawalls for better protection of coastal areas.

79 Overtopping flows can apply violent impact loads on seawalls and cause damage to coastal 80 infrastructures. In studies of wave overtopping loads, Oumeraci et al. (1993) classified wave-breaking 81 types and qualitatively linked slamming forces to wave shapes. Neelamani et al. (1999) established 82 empirical formulas for wave pressure prediction, emphasizing reflection and phase shifts. Cuomo et al. 83 (2010) conducted experiments within the VOWS (Violent Overtopping by Waves at Seawalls) 84 framework and observed discrepancies between empirical predictions and experimental data of wave 85 impact loads. In the combined experimental and numerical work by Hsiao and Lin (2010), the authors 86 found that the maximum dynamic net wave force on the seawall caused by solitary wave overtopping 87 typically corresponded to the peak surface elevations during overtopping. In the CPM simulation study 88 of Luo et al. (2019), it has been demonstrated that front slope angles significantly influence the forces 89 induced by overtopping flows. Qu et al. (2022) have shown that wind enhances wave propagation and 90 intensifies wave impact loads on a seawall. Recent research has examined the threats posed by wave 91 overtopping to pedestrians on coastal seawalls (Cao et al., 2021, Chen et al., 2021, Zhao et al., 2024), 92 focusing on the characteristics of overtopping flow depths and impact forces, as well as predictive 93 methodologies. Liu et al. (2023) investigated the dynamic response of viscoelastic floating covers under 94 wave overtopping conditions through Smoothed Particle Hydrodynamics simulations. While significant 95 progress has been made in understanding the dynamic processes of wave overtopping on seawalls, 96 critical questions remain regarding the behavior of overtopping characteristics and loads under extreme 97 wave conditions and their dependence on wave parameters.

98 This study investigates wave overtopping on a low-crested seawall under extreme wave conditions 99 through controlled wave flume experiments. A 1/16 scale model of a seawall section was constructed. 100 An increased water depth relative to the scaled on-site water depth was considered to simulate low-crest 101 scenarios, representing water level rises due to short-term wave setups. Extreme waves were generated

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using the focused wave theory with specified amplitudes and frequency bands. Comprehensive measurements were conducted, including wave elevations, deformed wave profiles, impact pressures and overtopping volumes. Detailed analyses were performed on the morphological features, energy evolutions, overtopping volumes and impact pressures of the wave overtopping flows, as well as their correlations with the water depth and wave parameters. The experimental methodology is presented in Section 2, results and discussion in Section 3, and conclusions in Section 4.

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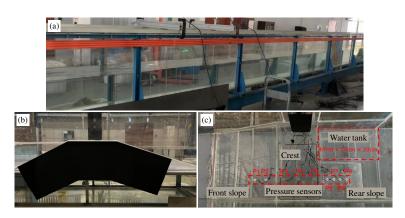
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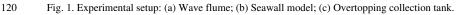
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109 2. Experimental methodology

110 2.1. Experimental model setup

The physical model experiments were conducted in a wave flume at Zhejiang University. The wave flume, 35 m in length, 0.6 m in width, and 0.8 m in depth, is equipped with a piston-type wave generator capable of generating both regular and irregular waves, as shown in Fig. 1(a). Unidirectional waves were considered in experiments and a scaled seawall with a uniform cross section (along the flume width direction) was deployed in the wave flume. The cross-section of the experimental seawall was modeled after typical coastal seawalls found along Zhejiang coastlines and a model scale ratio of $\lambda_1 = 1/16$ was adopted. The physical model comprised three main components:





(1) Slope bottom: representing the natural seabed, a gradient of 1:20 was employed. This slopewas located 10 m from the wave generator and extended 5 m in length.

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(2) Seawall structure: an impermeable seawall was constructed with a front slope of 1:3, a crest
height of 0.45 m, a crest width of 0.325 m, and a rear slope of 1:2 extending 0.4 m, as shown in Fig.
1(b).

127 (3) Overtopping collection system: positioned on the rear slope, a tank with dimensions of 0.4 m 128 $\times 0.2 \text{ m} \times 0.2 \text{ m}$ was used to collect overtopping water, as shown in Fig. 1(c).

To ensure a watertight assembly, silicone sealant was applied to fill gaps between the sloped bottom, seawall, flat bottom, and flume sidewalls. The experimental setup is illustrated in Fig. 2.

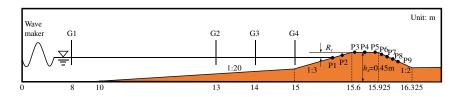


Fig. 2. Schematic view and key dimensions of the experimental setup of wave overtopping on a lowcrested seawall. The *x* coordinates of the pressure measurement points, i.e., P1 – P9, are $x_{P1} = 15.4$ m, $x_{P2} = 15.5$ m, $x_{P3} = 15.64$ m, $x_{P4} = 15.7625$ m, $x_{P5} = 15.885$ m, $x_{P6} = 15.985$ m, $x_{P7} = 16.045$ m, $x_{P8} = 16.105$ m, and $x_{P9} = 16.165$ m. The *x* coordinates of the wave elevation measurement points, i.e., G1 – G4, are $x_{G1} = 8$ m, $x_{G2} = 13$ m, $x_{G3} = 14$ m, and $x_{G4} = 15$ m.

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139 2.2. Measurement devices

140 The following instrumentation was deployed during the experiments:

141 (1) Wave gauges: Four KENEN capacitive wave gauges were installed at 8 m (G1), 13 m (G2), 14 142 m (G3), and 15 m (G4) from the wave generator to record free surface elevations. The wave gauges 143 have a full scale of -0.25 m to 0.25 m, a measurement resolution of 10^{-4} m, and a sampling frequency 144 of 100 Hz.

(2) Pressure sensors: nine piezoresistive pressure sensors were deployed to measure wave impact pressures on the seawall induced by overtopping flows. Two sensors were placed on the front slope, three on the crest, and four on the rear slope, with their precise locations detailed in Fig. 2. The pressure sensors have a measurement full scale of 5 kPa, a measurement resolution of 0.5% full scale, and a sampling frequency of 400 Hz.

(3) Overtopping volume measurement: the collected overtopping water was weighed and converted to volume, which was then divided by the tank's width (B = 0.2 m) to get the overtopping volume per unit width.

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(4) High-speed camera: An FR400 high-speed camera was adopted to capture the evolution of the
water surface morphology during wave overtopping, providing detailed visual data of wave-seawall
interactions. The camera has a sampling frequency of 100 Hz and a resolution of 2048 × 2048 pixels.

157 2.3. Wave generation methodology

158 In this study, extreme waves were modeled using the focused wave theory, which simulates the 159 dispersive focusing of a series of wave components. The wave elevation η is expressed as:

$$\eta(x,t) = \sum_{i=1}^{N} a_i \cos(k_i (x - x_f) - 2\pi f_i (t - t_f))$$
(1)

where, *N* is the number of wave components and N = 30 is adopted in the present study; a_i , k_i and f_i are the amplitude, wave number, and frequency of the *i*-th wave component, respectively. The terms x_f and t_f denote the focusing location and time of the wave group, respectively. The amplitude a_i of each wave component is determined by considering the JONSWAP (Joint North Sea Wave Project) wave spectrum (Goda, 2010), which reads:

$$S(f) = \beta_{\rm J} H_{\rm s}^2 f_{\rm p}^4 f^{-5} \exp(-1.25(f/f_{\rm p})^{-4}) \gamma^{\exp(-\frac{(f/f_{\rm p}-1)^2}{2\sigma^2})}$$
(2)

165 with

$$\beta_J = \frac{0.06238}{0.230 + 0.0336\gamma - 0.185(1.9 + \gamma)^{-1}} (1.094 - 0.01915\ln(\gamma))$$
(3)

166 and

$$\sigma = \begin{cases} 0.07 & \text{if } f \le f_p \\ 0.09 & \text{if } f > f_p \end{cases}$$
(4)

where f_p is the peak frequency; *Hs* is the significant wave height and is two times the amplitude of the focused wave at the focusing location (i.e., A_P); γ is the peak enhancement factor and is taken to be 3.3. With the wave spectrum density value for a wave component of frequency f_i , i.e., $S(f_i)$ and the amplitude of the focused wave at the focusing location (i.e., A_P), the amplitude of the *i*-th wave component can be calculated as:

$$a_{i} = A_{p} \frac{S(f_{i})\Delta f}{\sum_{i=1}^{N} S(f_{i})\Delta f}$$
(5)

172 where the angular frequency range Δf is defined as $\Delta f = \frac{f_{\text{max}} - f_{\text{min}}}{N}$, with f_{max} and f_{min} being the upper 173 and lower limits of the wave frequency band.

175 2.4. Experimental cases

176 The experimental parameters were designed to reflect realistic coastal conditions in Zhejiang 177 Province, China. By considering the experimental scale ratio, the first experimental water depth was set 178 at h = 0.40 m, representative of the typical coastal depths. The second depth was h = 0.45 m, which 179 took into account the water level rises due to storm surges and long-term sea-level rise. Correspondingly, 180 the crest elevations were $R_c = 0.05$ m and 0 m. The wave amplitudes (i.e., A_p) of 0.06 m, 0.07 m and 181 0.08 m were selected, along with frequencies (i.e., f_p) of 0.6 Hz \sim 0.8 Hz and 1.0 Hz. Considering the 182 randomness in the focusing location of extreme waves near the coast, seven focusing locations (i.e., x_f) 183 ranging from 12 m to 15 m with an interval of 0.5 m were investigated. With these, a total of 126 test 184 cases were established, as summarized in Table 1.

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Table 1. Summary of experimental cases

Parameters	Values					
<i>h</i> (m)	0.4, 0.45					
$R_{\rm c}$ (m)	0.05, 0					
$x_{\rm f}$ (m)	12, 12.5, 13, 13.5, 14, 14.5, 15					
$A_{\mathrm{p}}\left(\mathrm{m} ight)$	0.06, 0.07, 0.08					
$f_{\rm p}({\rm Hz})$	0.6	0.8	1.0			
$[f_{\min}, f_{\max}]$	[0.3, 0.9]	[0.5, 1.1]	[0.7, 1.3]			



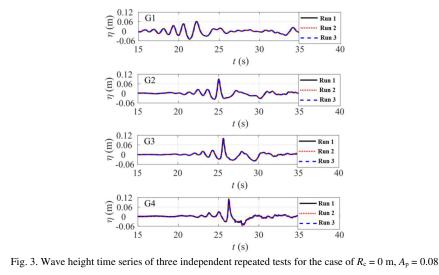
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m, $f_p = 0.6$ Hz and $x_f = 13$ m.

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193 2.5. Repeatability of experimental data

194 Experimental cases were repeated three times to ensure consistency. As an example, the wave 195 elevations and impact pressures for three repeated runs of the case characterized by $R_c = 0$ m, $A_p = 0.06$ m, $f_p = 0.6$ Hz and $x_f = 13$ m are shown in Figs. 3 and 4, respectively. The results demonstrate that the 196 197 wave heights and pressure time histories across the three independent repetitions nearly overlap, with a 198 relative error of less than 2%. The average overtopping volume measured across the three tests was 41.1 199 L/m, with an average relative error of just 0.49%. These findings highlight the good repeatability of the 200 wave generation system and the reliability of the measurement devices, confirming the accuracy of the 201 experimental data.



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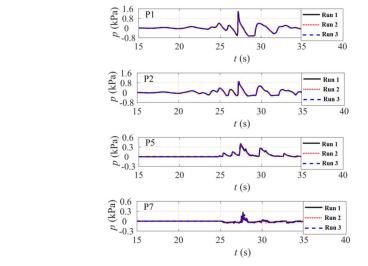


Fig. 4. Pressure time series of three independent repeated tests for the case of $R_c = 0$ m, $A_p = 0.08$ m, $f_p = 0.08$ m, $f_p = 0.6$ Hz and $x_f = 13$ m.

210 **3. Results and discussion**

211 3.1. Hydrodynamic characteristics of overtopping flows

To examine the hydrodynamic variations during the overtopping process of extreme waves on lowcrested seawalls, this section considers specific standard test conditions, including a spectral peak frequency $f_p = 0.6$ Hz, wave amplitude $A_p = 0.08$ m, focusing location $x_f = 13$ m, and seawall crest elevations $R_c = 0$ m and 0.05 m. These parameters were selected to facilitate a mechanistic analysis of hydrodynamic characteristics during overtopping under extreme wave conditions. The investigation

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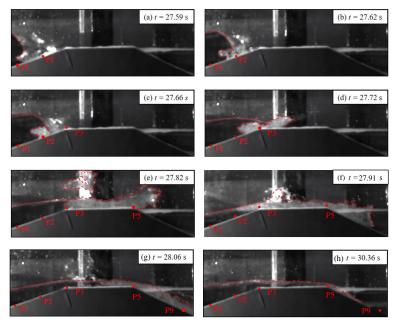
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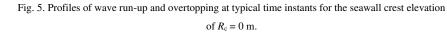
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AIP Publishing focuses on the overtopping patterns of extreme waves and the evolution of free surface elevations andthe corresponding wave energy.

219 3.1.1. Overtopping morphology evolution

220 The wave run-up and overtopping morphology for $R_c = 0$ m at representative time steps are shown 221 in Fig. 5. As the wave runs up along the sloping beach, the phenomenon of local water level dropping, 222 i.e., wave setdown, happens. Then, the leading edge of the primary wave crest steepens and plunges to 223 form wave breaking (Fig. 5a) and entraps some air (Fig. 5b). Subsequently, the plunging breaker 224 impacts on the front slope of the seawall, generating white splashes (Fig. 5c). As the wave continues its 225 run-up, the air pocket disintegrates entirely, generating sprays and a violent jet striking the seawall crest 226 (Fig. 5d), and the bubbly flows result in significant energy dissipations. The overtopping water further 227 generates an upward-impinging breaking wave near the crest's front edge, with the bubbly flow 228 transitioning into a rapid-moving jet across the rear crest edge (Fig. 5e). After that, a secondary wave 229 crest (hereafter termed the "secondary trailing crest") reaches the seawall's front slope without breaking 230 and causes more overtop flows Fig. 5f). Also note that Fig. 5(a), (c), (e), and (g) respectively correspond to the wave profiles when the pressure peaks at P1, P2, P3, and P5 occur. As can be seen that the 231 232 pressure peaks are primarily caused by breaking wave impacts.





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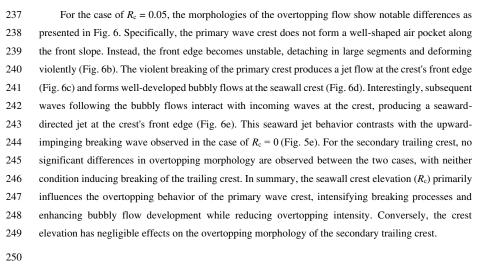
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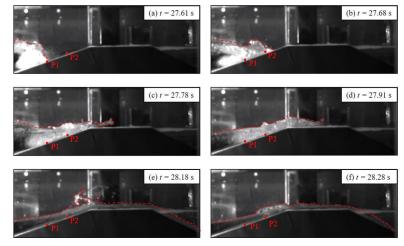


Fig. 6. Profiles of wave run-up and overtopping at typical time instants for the seawall crest elevation of $R_c = 0.05$ m.

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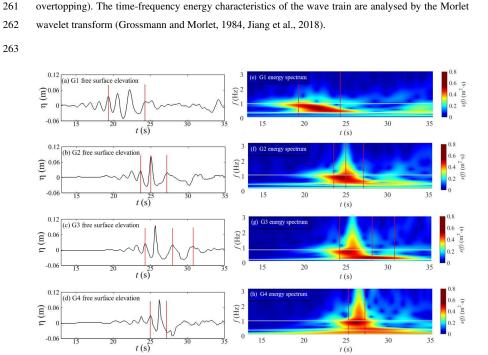
255 3.1.2. Spatio-temporal evolution of wave elevations and energy

In the process of wave run-up along the sloping seawall, wave shoaling and breaking occur. These are accompanied with wave energy transports and evolutions, which directly affect the wave impact

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loads applied on the seawall and the overtopping volume. Therefore, the spatiotemporal evolutions of

the wave elevations and their energy at typical locations are studied based on the representative case of $R_{\rm c} = 0$ m, $f_{\rm p} = 0.6$ Hz, $A_{\rm p} = 0.08$ m and $x_{\rm f} = 13$ m (this case is characterized by violent wave run-up and

Fig. 7. Time series of free surface elevations at G1-G4 (the left column) and the spatiotemporal energy evolutions (the right column).

267 At a location on the flat bed and 2 m away from the sloping beach toe (i.e., G1), relatively smaller 268 free surface elevations of the wave group are observed (Fig. 7a). The time-spectral contour of wave energy (Fig. 7e) shows that the wave energy focuses to some extent between t = 19.56 s to 23.69 s. The 269 270 peak wave energy of the wave train reaches 14.6 m² s upon the arrival of the primary wave crest at t =271 20.3 s. At the wave focusing location (i.e., G2), the wave group attains a large wave height (Fig. 7b) 272 with an almost symmetrical pattern for the primary crest. The energy spectrum contour (Fig. 7f) shows 273 that during t = 23.55 s to 25 s, shallower water depths amplify nonlinear effects, shifting energy from 274 the primary frequency to higher frequencies. The total wave group energy reaches 19.5 m²·s. After the 275 secondary wave crest, low-frequency wave components of certain energy pass through G2, and the 276 energy is slightly higher than that at G1 after the secondary wave crest passes.

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At G3 that is 1 m towards the coast, the primary wave crest becomes steeper and higher and less symmetric compared to that at G2 (Fig. 7c). Energy shifts from the primary frequency to higher frequencies during t = 24.29 s to 25.7 s. The time instants coincide with the arrival time of the secondary and primary wave crests (Fig. 7g). The shoaling effect at this location increases the amplitude and energy of the primary crest, peaking at 20.7 m²·s. After the secondary crest passes, waves with frequencies above 0.54 Hz propagate beyond G3, with energy concentrated between 0.3 Hz and 0.54 Hz.

284 At the seawall toe (i.e., G4), the shoaling effect intensifies further, which cause further increase of 285 the primary wave crest and pronounced asymmetry between the adjacent secondary crests at both sides 286 (Fig. 7d). Upon the primary crest's arrival (t = 26.28 s), the wave group energy reaches its maximum of 287 26.2 m^2 s (Fig. 7h). It can be seen from the energy spectrum contour that the wave energy shifts to both 288 the high- and low-frequency ranges during the running up process. This is related to the fact that the 289 reduced water depth causes the phase delay of the low-frequency long waves. Besides, high-frequency 290 wave components break on the seawall, releasing some wave energy, and some wave energy is reflected 291 back to the sea. These lead to an increase in the low-frequency wave energy at G4.

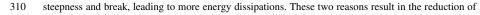
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293 3.2. Variation of wave energy evolutions with wave parameters and water depth

The wave energy evolution features during the wave run-up and overtopping process vary with the wave parameters (e.g., wave focusing location x_f and the spectral peak frequency f_p) and the water depth (which is related to the crest elevation R_c). These variations are critical to understanding the mechanisms of the overtopping behaviors induced by extreme waves. This section explores these influences based on the standard case (with the parameters of $R_c = 0$ m, $A_p = 0.08$ m, $f_p = 0.6$ Hz and $x_f = 13$ m) by analyzing the wave energy characteristics at locations G1 to G4.

300 3.2.1. Influences of wave focusing location

301 Fig. 8 (a) illustrates how the shift in wave focusing location (i.e., x_f) alters the wave group energy. 302 As the focusing location moves towards the shoreline, the low-frequency wave energy decreases 303 significantly and its duration shortens; the high-frequency energy also slightly decreases; the wave 304 energy concentrates more around the primary frequency. Specifically, when x_f shifts from 12 m to 15 m, the maximum energy of the wave group at G4 decreases from 26.2 m² to 24.7 m². These phenomena 305 306 are primarily attributed to two reasons. Firstly, with the increase in x_f , the energy focus of the wave train 307 is less developed at G4, and hence the wave height at G4 decreases (see Fig. 8 b). Secondly, the water 308 depth at the wave focusing location becomes shallower as $x_{\rm f}$ moves towards the shoreline and the 309 nonlinearity of the wave group is intensified. Hence, more wave components reach their maximum



311 the overtopping volume.

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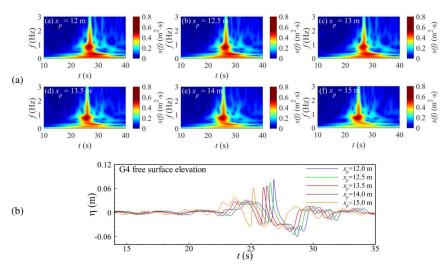


Fig. 8. Energy spectra and time histories of wave elevations at G4 in cases of different $x_{\rm f}$.

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315 3.2.2. Influences of spectral peak frequency

316 The spectral peak frequency (f_p) determines the energy distribution within the wave group and 317 influences the energy evolution along the wave propagation and run-up process. This section compares the energy spectra of wave elevations at locations G1-G4 for three different spectral peak frequencies, 318 319 i.e., $f_p = 0.6$ Hz, 0.8 Hz and 1.0 Hz. As shown in Fig. 9, the frequency ranges of waves with non-320 negligible energy for the three f_p values match the frequency ranges of the energy input from the wave 321 generator. As f_p increases, in general, the energy distributions of the waves at G1-G4 all shift towards 322 higher frequencies. More specifically, when $f_p = 0.6$ Hz the wave group contains more low-frequency wave energy particularly at location G4, whereas at $f_p = 0.8$ Hz and 1.0 Hz, the wave energy mainly 323 324 concentrates around the peak frequency, with only a small portion of the wave energy transferring to 325 the high- and low-frequency ranges. Besides, as the spectral peak frequency increases, the frequency 326 ranges of energy distributions narrow especially at G2-G4, implying the wave energy becomes more 327 concentrated. Moreover, the increase of f_p intensifies wave breaking, and hence more wave energy is 328 dissipated during the wave run-up process.

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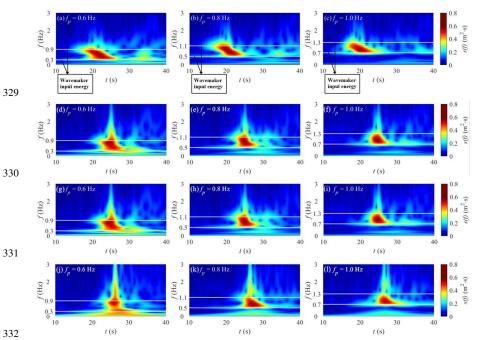


Fig. 9. Energy spectra of the wave elevation series at G1-G4 (the first to fourth rows, respectively) for different spectral peak frequencies f_p ($R_c = 0$ m, $A_p = 0.08$ m, $x_f = 13.0$).

336 3.2.3. Influences of seawall crest elevation

337 The water depth or crest elevation (R_c) significantly affects the wave energy evolution during the 338 wave run-up and overtopping. To investigate this, the wave elevations at G1-G4 for the cases of $R_c = 0$ 339 m and 0.05 m, $A_p = 0.08$ m, $f_p = 0.6$ Hz and $x_f = 13$ m are studied. As can be seen from Fig. 10, minor 340 differences exist in the wave energy spectra at G1 for $R_c = 0$ m and 0.05 m, because G1 is relatively far 341 from the wave focusing point and the nonlinear interactions among wave components are not intensive. At the wave focusing location G2, in the case of $R_c = 0.05$ m, more wave energy in the primary 342 343 frequency range shifts to the low and high-frequency ranges. This is attributed to the shallower water 344 depth, in which the nonlinear interactions among wave components and wave breaking are more 345 significant, increasing the wave energy in low frequencies. At G3, the high-frequency wave energy for 346 $R_{\rm c} = 0.05$ m is significantly reduced compared to $R_{\rm c} = 0$ m, while after the primary wave crest passes 347 the wave energy spectra for both seawall crests show little differences. At G4, wave breaking happens 348 in both cases; hence the low-frequency components are significant in the wave energy spectra.

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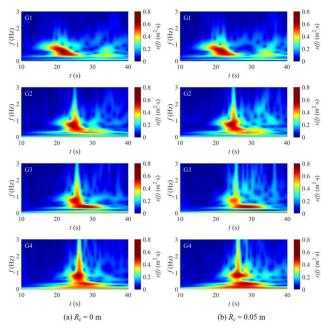


Fig. 10. Energy spectra of the wave elevation series at G1-G4 for different seawall crest elevations: (a) $R_c = 0$ m; (b) $R_c = 0.05$ m ($A_p = 0.08$ m, $f_p = 0.6$ Hz, $x_f = 13.0$).

354 3.3. Variation of overtopping volumes with wave parameters and water depth

This section examines the overtopping volume per unit width of extreme waves on a low-crested seawall (referred as "overtopping volume" hereafter) and analyzes its variations with the factors, such as the wave group amplitude A_p , wave focusing location x_f , spectral peak frequency f_p , and seawall crest elevation R_c .

359 3.3.1. Influences of wave amplitude

360 This section investigates the variation of wave overtopping volume with the focusing wave 361 amplitude A_p and wave focusing locations x_f under the condition of $R_c = 0$ m. As illustrated in

Fig. 11, for fixed R_c , f_p and x_f , the overtopping volume exhibits an almost linear increase with the focusing amplitude. For the three studied spectral peak frequencies, the rates of increase in overtopping volume when A_p rises from 0.06 m to 0.08 m are 41.1%, 37.0% and 52.2%, respectively, with the average value being 43.5%. The maximum overtopping volume observed in the studied cases is 40.50 L/m, occurring in the case of $R_c = 0$ m, $f_p = 0.6$ Hz, $A_p = 0.08$ m and $x_p = 12.0$ m. Conversely, the

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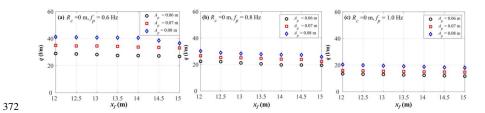
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368 and $x_p = 15.0$ m. In summary, these findings emphasize that wave overtopping volume increases sharply 369 with wave amplitude, underscoring the significant risks posed by extreme waves with large amplitudes 370 to seawall structural integrity and the potential for coastal flooding. 371

minimum overtopping volume, 12.12 L/m, is observed in the case of $R_c = 0$ m, $f_p = 0.6$ Hz, $A_p = 0.06$ m



373Fig. 11. Variations of wave overtopping volume per unit width (i.e., q) with focusing wave amplitude374 A_p and focusing location x_f .

376 3.3.2. Influences of wave focusing location

377The occurrence of extreme waves in real nearshore environments is inherently stochastic. This378section investigates the influence of the wave group focusing location (x_f) on overtopping volume to379understand the potential impacts of spatial randomness. Under the experimental conditions, overtopping380volumes per unit width were compared for seven focusing locations, ranging from $x_f = 12$ m to 15 m in3810.5 m intervals. As illustrated in

382 Fig. 11, for fixed R_c , A_p and f_p , extreme waves focusing at $x_f = 12$ m generate the highest 383 overtopping volume. As the wave focusing location shifts closer to the shoreline (i.e., x_f from 12 m to 15 m), the overtopping volume decreases, with reductions ranging from 0.18 L/m to 4.30 L/m. This 384 trend is attributed to the loss of wave group energy by the time the wave group reaches the seawall, 385 386 thereby diminishing overtopping intensity. It is worth noting that the influence of $x_{\rm f}$ on overtopping 387 volume is relatively minor compared to the effect of focusing wave amplitude. For example, at $f_p = 0.6$ Hz and $A_p = 0.06$ m, the overtopping volumes for $x_f = 12$ m and 15 m are 29.29 L/m and 27.07, 388 respectively, corresponding to a reduction of 7.6%. Similarly, the reduction rates for $A_p = 0.07$ m and 389 $A_{\rm p} = 0.08$ m are 5.2% and 11.1%, respectively. 390

391 3.3.3. Influences of spectral peak frequency

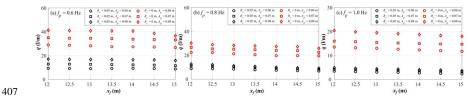
To investigate the influence of spectral peak frequency on overtopping volume, the overtopping volumes associated with extreme waves at $f_p = 0.6$ Hz, 0.80 Hz, and 1.0 Hz are compared. As shown in Fig. 12, for fixed R_c , A_p and x_p , extreme waves with $f_p = 0.6$ Hz produce the highest overtopping volume,

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followed by $f_p = 0.8$ Hz, with the lowest overtopping volume observed at $f_p = 1.0$ Hz. This trend can be 395 396 attributed to the fact that waves with lower spectral peak frequencies contain a larger proportion of low-397 frequency long waves, which exhibit relatively moderate nonlinear interactions during wave run-up and 398 retain their energy until breaking at the seawall crest. In contrast, waves with higher spectral peak 399 frequencies have a higher proportion of high-frequency components, which break prematurely on the seaward slope of the seawall, dissipating their energy and resulting in lower overtopping volumes. 400 401 Quantitatively, in cases of more intensive overtopping (i.e., $R_c = 0$ m and $A_p = 0.08$ m), when f_p increases 402 0.6 Hz to 1.0 Hz, the average overtopping volume across the seven studied $x_{\rm f}$ cases decreases by 20.81 403 L/m, corresponding to a reduction of 50%. For relatively moderate overtopping cases (i.e., $R_c = 0.05$ m 404 and $A_p = 0.06$ m), the average overtopping volume decreases by 6.27 L/m, representing a 66.7% 405 reduction.



408 Fig. 12. Variations of wave overtopping volume per unit width (i.e., q) with spectral peak frequencies 409 f_p and seawall crest elevation R_c .

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411 3.3.4. Influences of seawall crest elevation

412 Extreme waves are often accompanied by local sea level rises (e.g., those caused by storm surges), 413 which causes the seawalls to be low-crested ones. Hence, the influences of seawall crest elevation (R_c) 414 on overtopping are studied. As depicted in Fig. 12, as the seawall crest elevation reduces from 0.05 m 415 to 0 m, the overtopping volume increases from 9.73 L/m to 23.38 L/m, more than double. This 416 substantial increase is evident across all cases, demonstrating that seawall crest elevation is a critical 417 parameter influencing extreme wave overtopping volume. This underscores the wave overtopping risk 418 under extreme wave conditions with higher water levels, i.e., smaller R_c .

419 3.3.5. Summary of findings on overtopping volume

In summary, the above analyses suggest that the seawall crest elevation exerts the most significant
 influence on overtopping volume under extreme wave conditions, while the impacts of focusing wave
 amplitude, spectral peak frequency, and focusing location are comparatively less pronounced. Of these,

423 the wave focusing location has the least impact. Based on this, the wave overtopping volume across

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424 seven wave focusing locations are averaged for each combination of case parameters (i.e., R_c , f_p and A_p). 425 Table 2 presents the average overtopping volumes for two R_c , three A_p and three f_p . As f_p increases from 0.60 Hz to 1.0 Hz, the average overtopping volume decreases by 6.27 L/m (a reduction of two-thirds) 426 427 in the relatively mild overtopping condition (i.e., $R_c = 0.05$ m and $A_p = 0.06$ m) and by 20.81 L/m (a 428 reduction of one-half) in the relatively violent overtopping condition (i.e., $R_c = 0.0$ m and $A_p = 0.08$ m). 429 Additionally, Table 2 indicates that under wave conditions of low frequency and large amplitude, the 430 influence of reducing seawall crest elevation on overtopping volume is more pronounced. For example, 431 for the same focusing wave amplitude, the increase in overtopping volume for $f_p = 0.6$ Hz is 432 approximately twice that of $f_p = 1.0$ Hz; at the same spectral peak frequency, the increase in overtopping 433 volume for $A_p = 0.08$ m is about 1.4 times than that of $A_p = 0.06$ m.

434

Table 2. Average overtopping volume under different conditions

$R_{\rm c}$ (m) $A_{\rm p}$ (m)	Average overtopping volume(L/m)				
\mathbf{K}_{c} (III)	(m) A_p (m) $f_p = 0.6$ Hz 0.06 9.16 05 0.07 12.41 0.08 16.56 0.06 27.82	$f_{\rm p} = 0.8 \ {\rm Hz}$	$f_{\rm p} = 1.0 \; {\rm Hz}$		
	0.06	9.16	7.89	2.89	
0.05	0.07	12.41	9.84	3.76	
	0.08	16.56	10.66	4.21	
	0.06	27.82	20.74	12.61	
0	0.07	34.10	24.56	15.26	
	0.08	39.94	27.95	19.13	

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436 3.4. Hydrodynamic loading on the seawall

When extreme waves overtop a seawall, the resulting breaking phenomena generate significant
dynamic loads that challenge the structural integrity of the seawall. This section examines the wave
overtopping induced impact pressures on the front slope, crest, and rear slope of the seawall under
extreme wave conditions.

441 3.4.1. Time series of local pressures on the seawall

This section examines the time-varying characteristics of wave impact pressures on the seawall during overtopping, primarily based on the case with the most intense overtopping (i.e., $R_c = 0$ m, $A_p =$ 0.08 m, $f_p = 0.6$ Hz and $x_f = 13.5$ m). As shown in Fig. 13, the maximum overtopping pressures at all measurement locations, except for P6, are generated by the primary wave crest. At P6, the maximum pressure is generated by the secondary trailing crest. The time series of overtopping pressure reveals

overtopping process.

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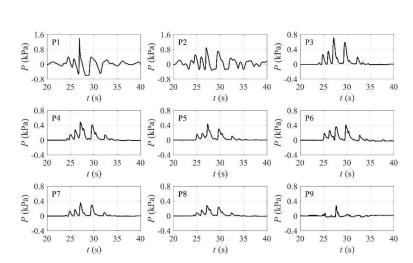
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that, apart from P9, multiple distinct pressure peaks are observed at P1 through P8 during the

wave crest (hereafter referred to as "primary crest pressure"), reaching 1.41 kPa. This value is

significantly higher than the pressures induced by secondary crests on either side ("secondary crest

pressure"), with the primary crest pressure being approximately four times larger. At P2 that is located

further upslope, the primary crest pressure decreases sharply to 0.89 kPa, while the impact pressures by

the secondary crests that have moderate values do not change significantly. Such general trends are

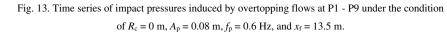
attributed to the behaviour of extreme waves as they run up along the seaward slope. During run-up,

the wave train continuously breaks (resulting in energy dissipation) and a portion of the wave's kinetic

energy is converted into gravitational potential energy, which lead to the impact pressures of

overtopping flows decreasing progressively along the seaward slope.

On the seaward slope of the seawall, P1 records the highest overtopping pressure from the primary



464 On the seawall crest, the time series of overtopping pressures at P3 to P5 shows distinct pressure 465 peaks. These peaks are generated by the primary wave crest, as well as the secondary and tertiary crests 466 on either side. This observation suggests that the overtopping pressure on the crest is primarily caused 467 by the primary wave crest and the closely following secondary crests, consistent with the finding that 468 wave overtopping in low-crested seawalls under extreme waves is predominantly caused by the primary and secondary trailing crests. Intense wave breaking on the crest leads to significant energy dissipation, 469

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470 resulting in a progressive decrease in overtopping pressure along the crest. Accordingly, the 471 abovementioned pressure peaks show a declining trend as they propagate along the crest.

472 On the rear slope of the seawall, the maximum overtopping pressure at P6 is generated by the 473 secondary trailing crest. This is likely because, as the primary crest reaches the rear edge of the seawall 474 crest, it forms a jet that fully breaks before reaching P6 (see Fig. 14). Consequently, P6 is not fully 475 impacted by the primary crest. Instead, the secondary trailing crest flows down the rear slope, directly 476 impacting P6 and resulting in slightly higher pressure than that produced by the primary crest. At P7 477 and P8, both the primary and secondary trailing crests contribute to the overtopping pressure, while P9 478 is influenced only by the primary crest. A comparison of the overtopping pressure time series from P6 479 to P9 shows that the maximum pressure on the rear slope decreases progressively along the slope. This 480 decrease is related to the energy dissipation caused by wave breaking, which reduces its impact pressure. 481 With wave propagation, the primary crest pressure exhibits only a modest decrease, while the secondary 482 crest pressure drops significantly (eventually reaching negligible levels). This indicates that the 483 secondary trailing crest loses energy more rapidly than the primary crest as it propagates along the rear 484 slope of the seawall.

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487 Fig. 14. Overtopping pattern of the overtopping jet flow just as it contacts the rear slope at position P6 488 during the extreme wave overtopping process.

490 3.4.2. Variation of the maximum impact pressures by overtopping

491 (1) Influences of wave amplitude and wave focusing location

492 Based on the experimental cases of relatively intense wave overtopping (i.e., $R_c = 0$ m and $f_p = 0.6$ Hz), the influences of focusing wave amplitude and wave focusing location on the peak impact pressure 493 494 are analysed, with the results presented in Fig. 15. Generally, for a fixed wave focusing location, the 495 peak impact pressure on the seawall increases linearly with focusing amplitude. Conversely, for a fixed 496 focusing amplitude, the peak pressure decreases as $x_{\rm f}$ shifts towards the shoreline. These trends align 497 with the observed variation of wave overtopping volume as a function of wave amplitude and wave 498 focusing location. Among the cases studied, the highest impact pressure occurs at P1, reaching 1.77 499 kPa, indicating that P1 is subjected to the most violent wave impacts.

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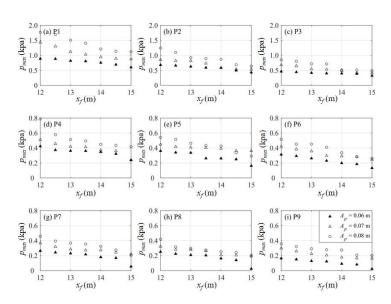


Fig. 15. Variations of peak slamming pressures at P1 to P9 with focusing amplitude A_p and focusing location x_p for the case of $R_c = 0$ m and $f_p = 0.6$ Hz.

505 To provide a more detailed evaluation, cases with fixed R_c , f_p and x_f are selected as illustrative 506 examples. Using the peak impact pressure induced by an extreme wave with a focusing amplitude of $A_p = 0.06$ m as the baseline, the rates of increase in impact pressures at P1 – P9 as A_p increases from 507 508 0.06 m to 0.08 m are calculated and summarised in Table 3. The results reveal that on the seaward slope 509 of the seawall, the impact pressure at P1 increases by 80.2%, which exceeds the 55.0% increase at P2, 510 indicating that impact pressures at P1 are more sensitive to variations in wave amplitude. Notably, at 511 P1, where the most violent impacts occur, the peak impact pressure exhibits a 99.1% increase as Aprises 512 from 0.06 m to 0.08 m. On the seawall crest, the increase rates at P3, P4, and P5 are comparable, 513 measuring 61.3%, 51.0%, and 54.7%, respectively. On the rear slope of the seawall, the rate of pressure 514 increase generally becomes larger as the measurement location moves towards the shoreline, with the 515 growth rates exceeding 100% in some cases.

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Measurement	Rates of pressure peak increase when $A_{\rm p}$ increases from 0.06 m to 0.08 m (%)						
locations	$x_{\rm f} = 12$	$x_{\rm f} = 12.5$	$x_{\rm f} = 13$	$x_{\rm f} = 13.5$	$x_{\rm f} = 14$	$x_{\rm f} = 14.5$	$x_{\rm f} = 15$
P1	99.05	92.64	83.83	74.59	60.24	64.62	86.67
P2	83.47	66.57	47.69	51.71	49.14	38.50	48.05
P3	84.79	82.06	72.51	78.33	28.58	32.79	49.77
P4	85.85	55.09	41.47	37.07	29.46	34.77	72.93
P5	50.00	51.85	37.51	64.71	63.40	35.22	80.08
P6	65.67	54.45	72.36	78.13	69.18	52.59	96.97
P7	73.86	62.42	59.92	63.27	81.32	63.03	273.75
P8	67.13	41.89	41.91	34.60	54.48	72.22	727.68
P9	118.36	115.10	126.39	126.82	200.06	152.52	783.94

Table 3. Increase rates of peak impact pressures at P1 to P9 for the cases of $R_c = 0$ m and $f_p = 0.6$ Hz 21 as A_p increases from 0.06 m to 0.08 m

522

523 (2) Influences of spectral peak frequency

524 Based on the cases with $R_c = 0$ m and $A_p = 0.08$ m, the influences of the spectral peak frequency 525 on peak impact pressures are investigated. Fig. 16 shows that the peak impact pressures at P1 to P9 rise as $f_{\rm P}$ decreases, consistent with the variation trend of overtopping volume. Analyzing the peak pressures 526 527 at P1 to P9 as f_p decreases from 1.0 Hz to 0.6 Hz, it is found that the changes of peak pressures with x_f 528 are within 1.23% to 4.87%, being not significant. Accordingly, the peak impact pressures among the 529 cases with different $x_{\rm f}$ are averaged. Based on the averaged values, the increase rates of the peak impact 530 pressures when fp changes from 1.0 Hz to 0.6 Hz are analyzed, as shown in Table 4. It can be seen that 531 the peak impact pressure increases more significantly on the seaward slope of the seawall as compared 532 to that on the seawall crest and rear slope. Notably, P1 exhibits the highest overtopping pressure and is 533 the most sensitive to the change of f_p . The peak pressure at P3 (located near the front edge of the crest) 534 is also evidently affected by f_p , while the peak pressures at locations closer to the shoreline (i.e., P4 to 535 P9) are less sensitive to f_p .

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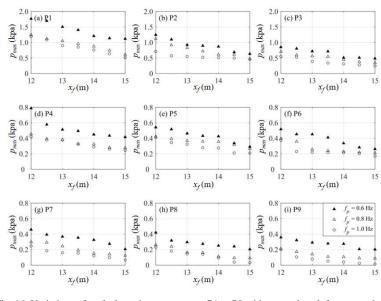
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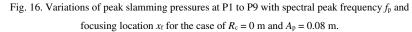
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541Table 4. Increase rates of peak impact pressures at P1-P9 for the cases of $R_c = 0$ m and $A_p = 0.08$ m542(the values of different x_f are averaged) as f_p changes from 1.0 Hz to 0.6 Hz

Measurement locations	P1	P2	P3	P4	Р5	P6	P7	P8	Р9
Increase rates of impact pressure (%)	55.6	36.1	27.8	19.6	13.6	15.4	19.3	16.3	20.2

544 (3) Influences of seawall crest elevation

545 Using cases with $A_p = 0.08$ m and $f_p = 0.6$ Hz, the variation of peak impact pressures on the seawall 546 with crest elevation R_c is examined. As shown in Fig. 17, the peak impact pressures at points P₁ to P₉ 547 increase significantly as R_c decreases. This occurs because a lower seawall crest elevation allows 548 extreme waves under similar conditions to carry more water during overtopping, resulting in higher 549 kinetic energy and hence consequently a substantial rise in slamming pressures on the seawall. For quantitative analysis, the peak impact pressures at each location are averaged for cases with the same 550 551 $R_{\rm c}$ and varying $x_{\rm f}$. The results indicate that when $R_{\rm c}$ decreases from 0.05 m to 0 m, the peak pressures at 552 P1 and P2 increase by 0.49 kPa and 0.48 kPa, respectively. On the seawall crest, the increases are

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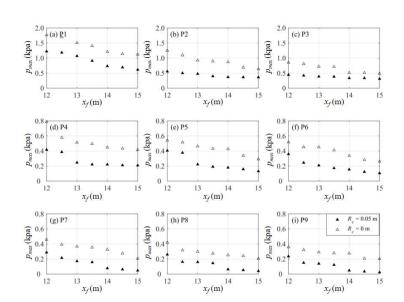
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566 567 affected by changes in crest elevation.

more sensitive to wave parameters and water depth.



relatively smaller, measuring 0.29 kPa, 0.25 kPa and 0.19 kPa at P3, P4 and P5, respectively. On the

rear slope of the seawall, the rise in peak pressures due to the reduction in R_c becomes even less pronounced. These findings suggest that impact pressures at locations closer to the sea are more strongly

In summary, the peak impact pressures on a seawall induced by overtopping flows increase with

greater focusing amplitude, lower spectral peak frequency, and reduced seawall crest elevation, while

they decrease as the wave focusing location (x_f) shifts towards the shoreline. These trends are consistent

with those observed for overtopping volume in relation to these parameters. Additionally, impact pressures at locations on the seaward slope and near the front edge of the seawall crest are generally

Fig. 17. Variations of peak slamming pressures at P1 to P9 with seawall crest elevation R_c and focusing location x_f for the case of $A_p = 0.08$ m and $f_p = 0.6$ Hz.

568 4. Conclusions and perspectives

569 This study has conducted well-controlled wave flume experiments to investigate the overtopping 570 mechanisms of extreme waves on low-crested seawalls. Extreme waves were modeled using the linear 571 focusing of a wave train. The experiments measured wave elevations at typical locations, wave profiles 572 during overtopping, wave impact pressures on the seaward slope, crest and rear slope of the seawall,

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and overtopping volume per unit width. The morphological and dynamic features of the overtoppingflows have been analyzed. The key findings are summarized below:

(1) Wave breaking and energy dissipation: As the spectral peak frequency and seawall crest elevation increase, the primary wave crest tends to break earlier on the seawall's front slope. This results in intensified breaking and significant energy dissipation. Additionally, when the wave focusing location of the wave train shifts closer to the shoreline, the energy reaching the seawall to decreases, leading to reduced overtopping volumes and wave impacts.

(2) Influences of wave parameters on overtopping volume: Larger wave heights, lower spectral
peak frequencies, reduced seawall crest elevations (i.e., larger water levels) and seaward shifts in wave
focusing location result in more overtopping volume. Quantitatively:

- i) The maximum overtopping volume in the studied cases reaches 40.50 L/m, occurring when R_c = 0 m, f_p = 0.6 Hz, A_p = 0.08 m and x_p = 12.0 m. The minimum overtopping volume is 12.12 L/m, occurring when R_c = 0 m, f_p = 0.6 Hz, A_p = 0.06 m and x_p = 15.0 m.
- ii) Reducing the seawall crest elevation from 0.05 m to 0 m results in a substantial increase in overtopping volume, from 9.73 L/m to 23.38 L/m, more than doubling the overtopping discharge. Reducing the spectral peak frequency from 1.0 Hz to 0.6 Hz increases the overtopping volume from 6.27 L/m to 20.81 L/m. A shoreward shift in the wave focusing location (i.e., *x*_f from 12 m to 15 m) results in slight reductions in overtopping volume, ranging from 0.18 L/m to 4.30 L/m.

iii) These observations reveal a hierarchy of influences on overtopping volume, with seawall crest
 elevation exerting the strongest effect, followed by spectral peak frequency, focusing
 amplitude, and focusing location.

(3) Pressure distributions and variations: The impact pressures by overtopping flows of different wave parameters indicate that the peak impact pressure occurs at P1 with a value of 1.77 kPa, which corresponds to the case of $A_p = 0.08$ m, $R_c = 0$ m, $f_p = 0.6$ Hz and $x_f = 12.0$ m. At P1 where violent impacts happen, the peak impact pressure increases by 99.1% when the focusing amplitude A_p increases from 0.06 m to 0.08 m.

The insights gained from this study have significant implications for coastal engineering,
 particularly in the design and maintenance of low-crested seawalls to mitigate risks posed by extreme
 wave events. Key applications include:

i) Improved seawall design: the quantification of overtopping and pressure distribution enables
 engineers to refine seawall geometries, such as crest height and slope gradients, to balance
 safety and construction costs.

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- ii) Risk assessment and adaptation: the findings provide theoretical bases for assessing the vulnerability of coastal infrastructure under varying wave conditions, including scenarios exacerbated by climate change and sea-level rise.
- iii) Wave energy management: understanding the spatiotemporal evolution of wave energy facilitates the development of energy-dissipative devices and materials to enhance coastal resilience.

612 While this study provides a comprehensive understanding of overtopping dynamics, several 613 avenues for further research are identified to extend its scope, such as the influence of complex 614 bathymetry (future studies could explore how variations in seabed topography affect wave-seawall 615 interactions and overtopping behaviours), the impact of multi-directional waves (investigating the 616 effects of oblique or multi-directional waves would offer a more realistic representation of overtopping 617 dynamics), the material and structural innovations (research into advanced materials and novel 618 structural designs, such as hybrid seawalls or energy-absorbing components, could further enhance the 619 effectiveness of low-crested seawalls). By addressing these research directions, the understanding of 620 overtopping mechanisms and their applications can be expanded, ultimately contributing to more 621 resilient and sustainable coastal defence systems.

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